

New Understandings of the atmospheric turbulent mixing



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What I learned from BOREAS

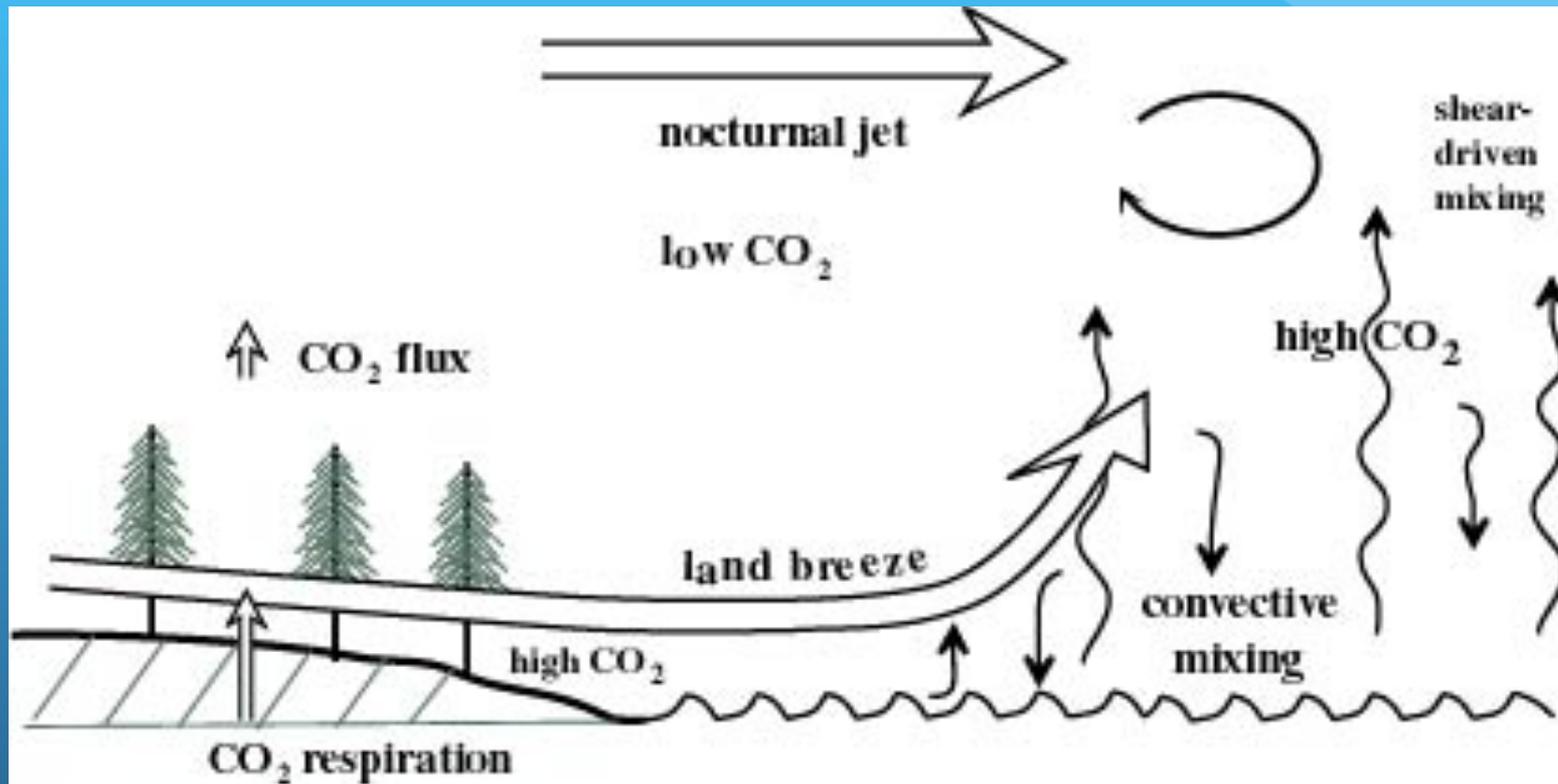
- **The first CO₂ advection experiment (nighttime CO₂ chimney)**
- **skin temperature vs. aerodynamic temperature**

New Understandings of the atmospheric turbulent mixing

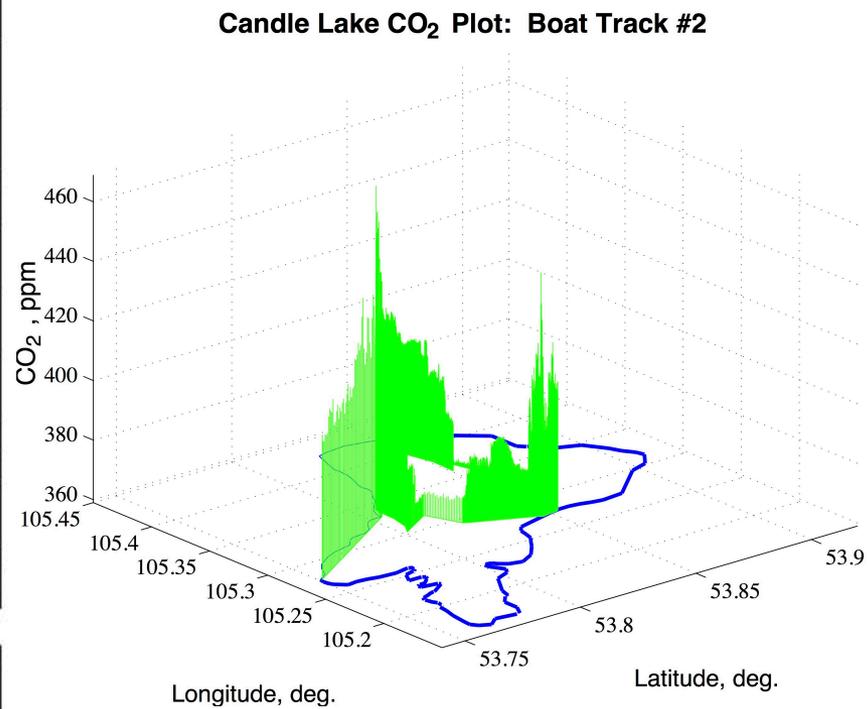
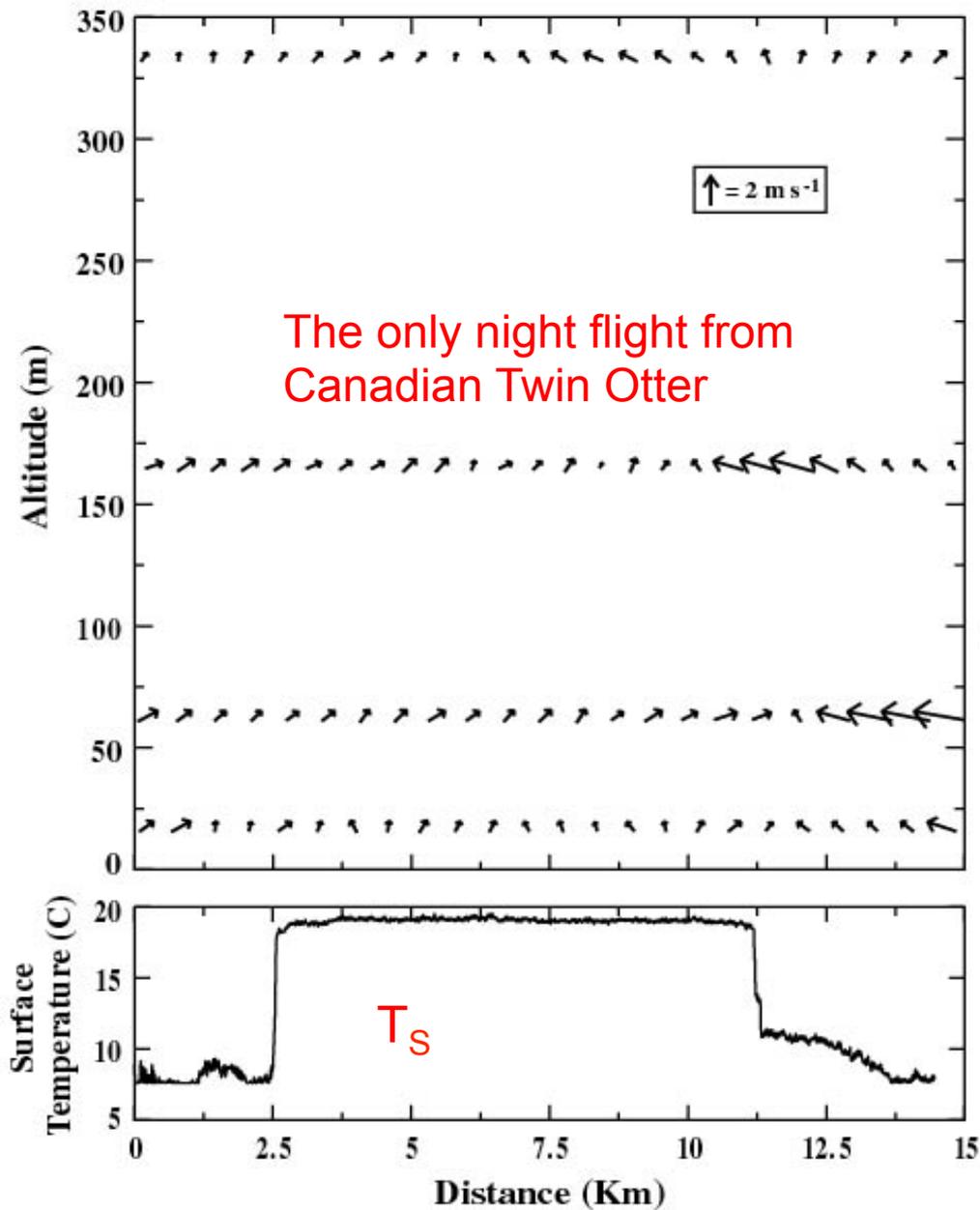
- **Molecular diffusion vs. turbulent mixing**
- **MOST vs HOST**

1. CO₂ advection

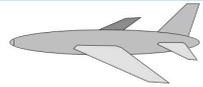
Sun, Desjardins, Mahrt, MacPherson (1998, JGR)



Marc Aubinet just won the award for outstanding achievement in biometeorology “for significant contributions,... to apply the eddy covariance method to atmosphere-biosphere interactions and to the problem of advection”



2. Surface skin temperature vs. aerodynamic temperature



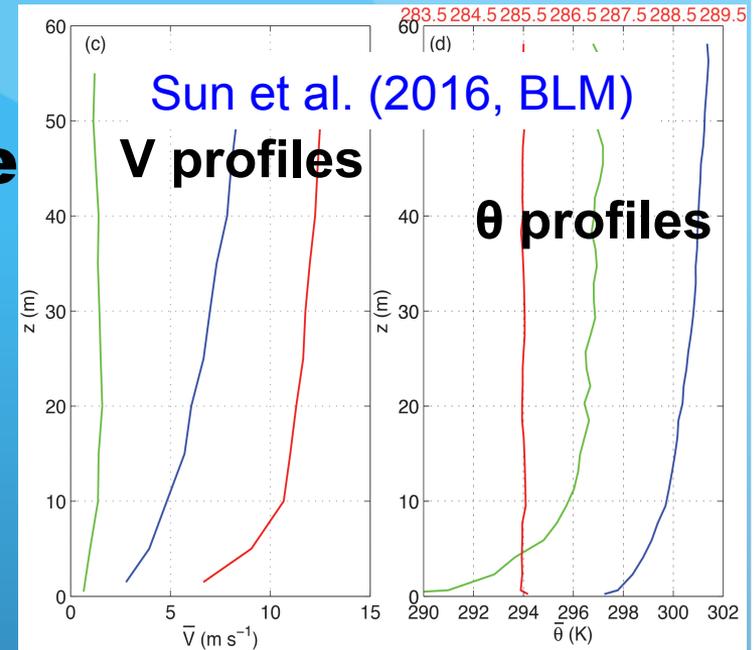
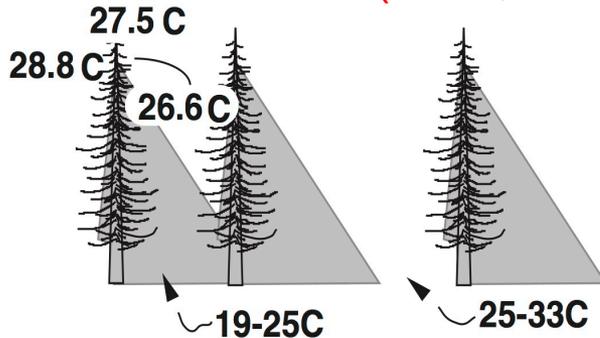
$$\overline{T_{air}} = 25\text{ C}$$

$$\overline{T_{sfc}} = 24\text{ C}$$

↑ observed
heat
flux

gradient
heat
flux ↓

Sun and Mahrt (1995, BLM)



$$H = C_h V (T_a - T_0) \quad \checkmark$$

$$H = C_h V (T_a - T_s) \quad \times$$

Using the skin temperature (T_s) in parameterization of H can lead to the run-away cooling

Molecular diffusion vs. turbulent mixing (1)

large scale

fast

small scale

slow

turbulent mixing (MOST-
Monin-Obukhov similarity theory)

molecular diffusion

Molecular diffusion vs. turbulent mixing (2)

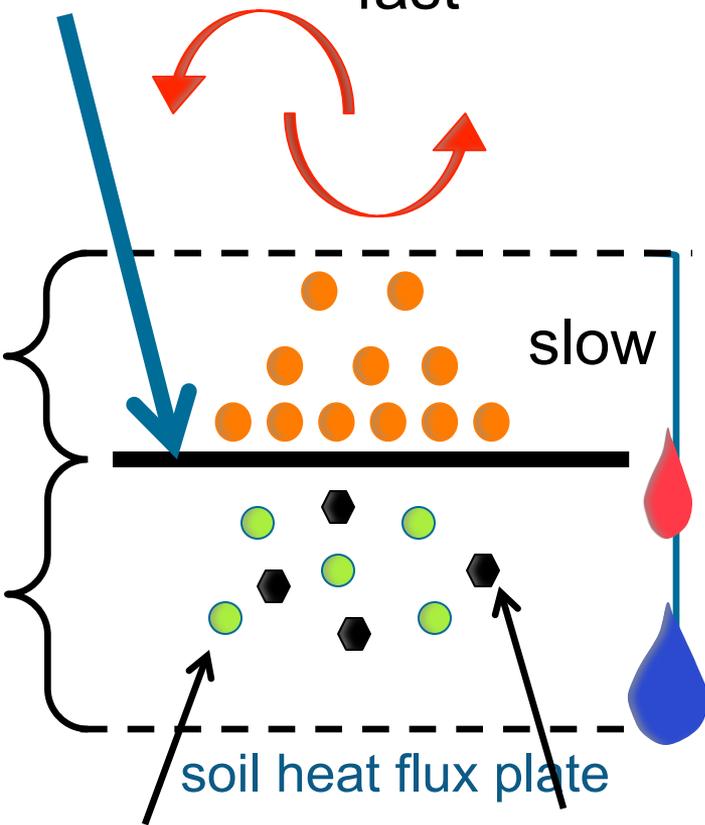
surface energy imbalance

sonic anemometer

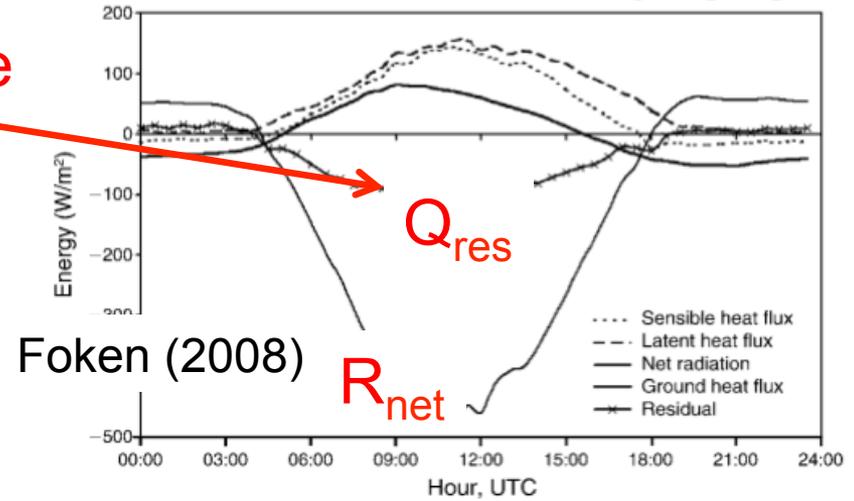


solar radiation

fast



water $\rho_w > \rho_s$
 $C_{pw} > C_{ps}$ soil



total energy conservation

$$c_p \frac{dT}{dt} + \frac{dE_k}{dt} = Q + v \frac{dp}{dx} + \epsilon_T$$

momentum conservation

$$c_p \frac{dT}{dt} + \epsilon = Q$$

Surface energy balance involves the thermal energy balance associated with radiation and molecular diffusion only

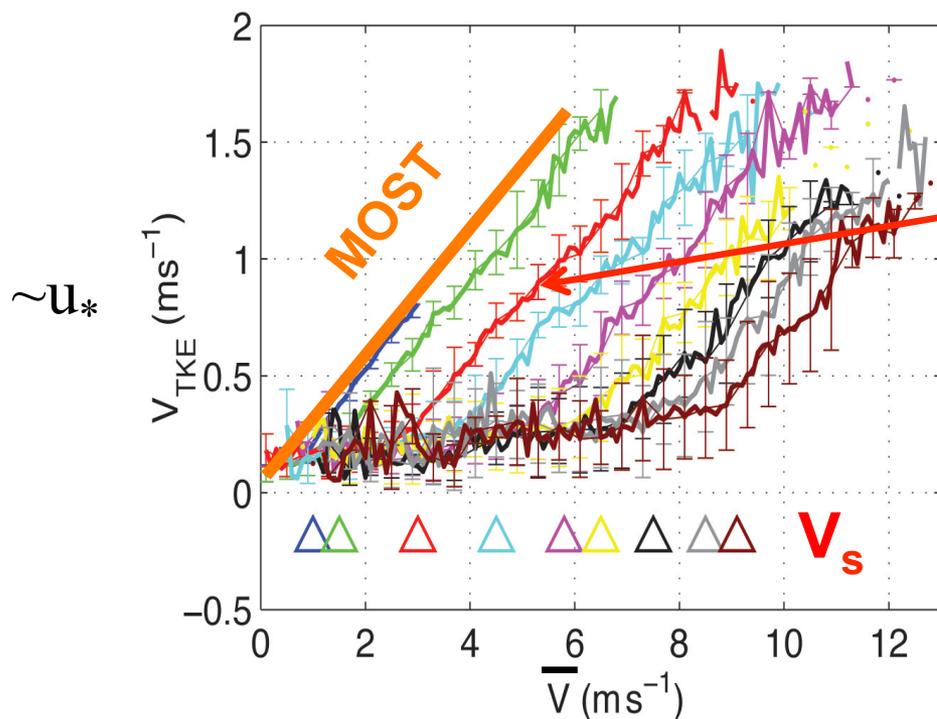
turbulent mixing near the surface

$$u_* = \overline{|w'V'|}^{1/2} \sim \delta V / \delta z \begin{cases} \delta z \rightarrow 0, \partial V / \partial z \\ \delta z = [0, z], \left[\frac{\partial V}{\partial z}, \frac{V}{z} \right] \end{cases}$$

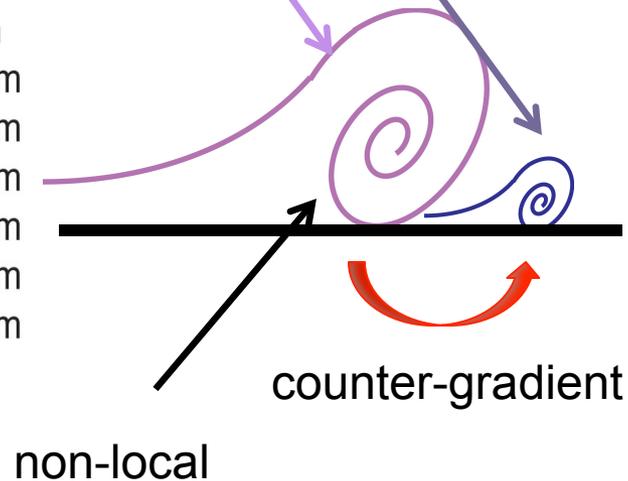
MOST $u_* \sim V(z) f(Ri, z)$

HOST $u_* \sim \alpha(z) + \beta(z)V(z)$

HOckey-Stick Transition (HOST)



- 0.5m
- 1.5m
- 5m
- 10m
- 20m
- 30m
- 40m
- 50m
- 55m



HOST hypothesis:

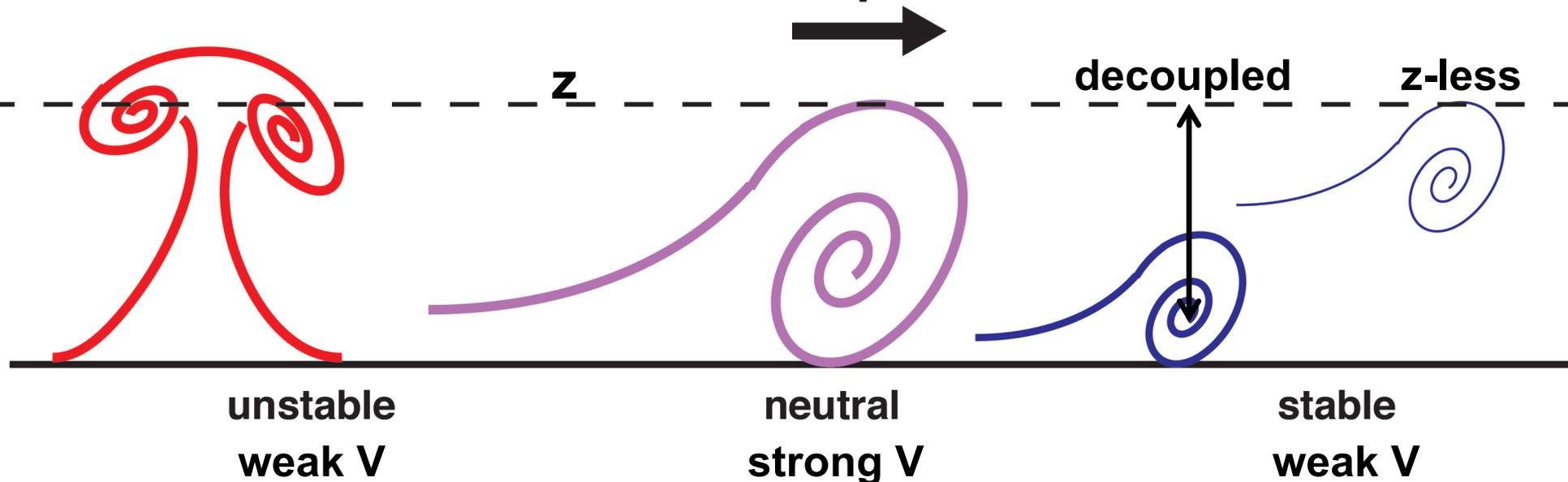
1. turbulence generation is on a finite scale, δz
2. turbulence intensity is determined by the TKE and TPE partition as well as turbulence generation in the δz layer

buoyancy generation

weak V
strong turbulence
 $\delta z \rightarrow z$

$\delta V / \delta z$ not $\partial V / \partial z$
strong V
strong turbulence
 $\delta z = z$

weak V
weak turbulence
 $\delta z < z$



summary

- **Need to understand aerodynamics (turbulent mixing, air-land interaction) in order to scale up from leaf to regions.**
- **New understandings of turbulent mixing (large coherent eddies controlled by the energy conservation) will shed light on how to keep track trace gases in the atmosphere-ecosystem interaction.**